

Interpreting GEMS geostationary satellite observations of the diurnal variation of nitrogen dioxide (NO₂) over East Asia

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Abstract. Nitrogen oxide radicals (NO_x ≡ NO + NO₂) emitted by fuel combustion are important precursors of ozone and particulate matter pollution, and NO₂ itself is harmful to public health. The Geostationary Environment Monitoring Spectrometer (GEMS), launched in space in 2020, now provides hourly daytime observations of NO₂ columns over East Asia. This diurnal variation offers unique information on the emission and chemistry of NO_x, but it needs to be carefully interpreted. Here we investigate the drivers of the diurnal variation of NO₂ observed by GEMS during winter and summer over Beijing and Seoul. We place the GEMS observations in the context of ground-based column observations (Pandora instruments) and GEOS-Chem chemical transport model simulations. We find good agreement between the diurnal variations of NO₂ columns in GEMS, Pandora, and GEOS-Chem, and we use GEOS-Chem to interpret these variations. NO_x emissions are four times higher in the daytime than at night, driving an accumulation of NO₂ over the course of the day, offset by losses from chemistry and transport (horizontal flux divergence). For the urban core, where the Pandora instruments are located, we find that NO₂ in winter increases throughout the day due to high daytime emissions and increasing NO₂/NO_x ratio from

40 entrainment of ozone, partly balanced by loss from transport and with negligible role of chemistry. In summer, by contrast, chemical loss combined with transport drives a minimum in the NO_2 column at 13-14 local time. Segregation of the GEMS data by wind speed further demonstrates the effect of transport, with NO_2 in winter accumulating throughout the day at low winds but flat at high winds. The effect of transport can be minimized in summer by spatially averaging observations over the broader metropolitan scale, under
45 which conditions the diurnal variation of NO_2 reflects a dynamic balance between emission and chemical loss.

1. Introduction

The Geostationary Environment Monitoring Spectrometer (GEMS) satellite instrument was launched in February 2020 by the National Institute of Environmental Research (NIER) to observe air quality over
50 East Asia. GEMS is the first geostationary instrument directed at air quality and provides hourly column measurements of several gases including nitrogen dioxide (NO_2) (J. Kim et al., 2020). NO_2 is part of the nitrogen oxides ($\text{NO}_x \equiv \text{NO} + \text{NO}_2$) radical family, which is emitted by fuel combustion and whose chemistry plays a critical role in driving ozone (O_3) and fine particulate matter ($\text{PM}_{2.5}$) formation. NO_2 itself is of concern as an air pollutant. Loss of NO_x is by atmospheric oxidation by the hydroxyl radical
55 (OH) and ozone, resulting in a lifetime of a few hours in summer and about a day in winter (Shah et al., 2020). The diurnal cycle of NO_2 measured from geostationary orbit offers unique information on the emission, chemistry, and transport of NO_x . Here we interpret the GEMS observations with the GEOS-Chem chemical transport model (CTM) to better understand the processes controlling this diurnal cycle.

Several studies have examined the diurnal variation of NO_2 in urban air using surface
60 concentrations from air quality networks. The data typically exhibit bimodal maxima in the morning around 7-9 local time (LT) and in the evening around 19-21 LT, including over Beijing and Seoul (Cheng et al., 2018; H. Kim et al., 2020). This has been commonly attributed to high NO_x emission during morning and evening rush hours (Kendrick et al., 2015; Cheng et al., 2018), but urban NO_x emission inventories show little variation during daytime (Miao et al., 2020). Moutinho et al. (2020) found that the morning and

65 evening NO₂ maxima could be driven by shallow mixing depths, in contrast to the middle of the day where
surface heating drives deep mixing and defines the depth of the planetary boundary layer (PBL) in daily
contact with the surface. The PBL depth extends typically to 1-3 km altitude.

Ground-based measurements of NO₂ columns are available from the Pandora Sun-staring
spectrometer instrument network used for validating satellite observations (Herman et al., 2009; Kanaya et
70 al., 2014; Judd et al., 2020; Verhoelst et al., 2021). Column measurements integrate concentrations from
the surface to the top of the atmosphere and are therefore not directly sensitive to mixing depth. The
Pandora network consists mainly of urban sites, where the NO₂ column and its variability are mainly within
the PBL (Yang et al., 2023). The Pandora data from Seoul tend to show an increasing trend in the early
morning followed by flat concentrations over the rest of the daytime, with less diurnal variation than NO₂
75 concentrations in surface air (Crawford et al., 2021). Nearby sites can show different diurnal variations,
pointing to a major role of local transport in driving this variation (Chang et al., 2022; S. Kim et al., 2023).

Satellite observations of NO₂ from polar sun-synchronous low-earth orbit (LEO) have been made
since 1995 starting with the GOME instrument (Martin et al., 2002) but observe by design at a single time
of day. Several studies have combined observations from the SCIAMACHY or GOME-2 instruments
80 observing in the morning at 9-10 LT and the OMI instrument observing in the afternoon at 13-14 LT to get
some information on NO₂ diurnal variation. Boersma et al. (2008) found decreases from morning to
afternoon over urban regions that they attributed to photochemical loss, and increases from morning to
afternoon over tropical biomass burning regions that they attributed to a midday maximum in emissions.
Boersma et al. (2009) found that the urban morning-to-afternoon decrease was largest in summer and
85 absent in winter. Penn and Holloway (2020) found that NO₂ column ratios between morning and afternoon
were lower than surface NO₂ concentration ratios, as would be expected from deeper vertical mixing in the
afternoon. Ghude et al. (2020) found an important role for transport in driving morning-to-afternoon
variations in NO₂ columns over urban India.

Here we analyze and compare the NO₂ diurnal cycles observed by GEMS over the Seoul and
 90 Beijing metropolitan areas in winter and summer. We compare to the diurnal cycles observed by Pandora in
 the urban cores and to simulations with the GEOS-Chem CTM. We use GEOS-Chem to separate and
 quantify the roles of emission, chemistry, and transport in driving the NO₂ diurnal cycles observed from
 GEMS over different spatial scales. This work provides a basis for more quantitative application of GEMS
 observations as top-down information on NO_x emissions, and more generally for interpreting the diurnal
 95 cycle of NO₂ from geostationary orbit with application to the TEMPO instrument over North America
 launched in April 2023 (Zoogman et al., 2017) and the Sentinel-4 instrument over Europe to be launched in
 2024 (Gulde et al., 2017).

2. Observations and model

2.1 GEMS data

100 GEMS is an ultraviolet-visible instrument measuring back-scattered solar spectra at 300 – 500 nm
 (J. Kim et al., 2020). It was launched in February 2020 in geostationary orbit at a longitude of 128.25°E. Its
 native pixel resolution is $3.5 \times 8 \text{ km}^2$ at Seoul, with variations in Level 2 product resolutions depending on
 the pixel binning configurations. We use hourly total NO₂ slant column density from the GEMS L2 NO₂
 version 2.0 product at $3.5 \times 8 \text{ km}^2$ resolution for December-February (DJF) 2021/22 and June-August
 105 (JJA) 2022 (NIER, 2023). The GEMS NO₂ algorithm uses differential optical absorption spectroscopy
 (DOAS) to fit back-scattered solar spectra within the 432 – 450 nm range (J. Kim et al., 2020). This yields
 the slant column density along the light path (L2 data). We use all GEMS L2 NO₂ version 2.0 data that pass
 algorithm quality flag ≤ 112 , final algorithm flag ≤ 1 , solar zenith angle (SZA) $< 70^\circ$, viewing zenith angle
 (VZA) $< 70^\circ$, and cloud fraction < 0.3 (Lee et al., 2020).

110 The vertical column density of NO₂ is obtained by dividing the slant column density by an air mass
 factor (AMF) characterizing the photon path from the Sun down through the atmosphere and back up to the
 instrument. The AMF depends on the viewing geometry and on the scattering properties of the atmosphere:

$$\text{AMF} = \text{AMF}_G \int_0^\infty w(z) S(z) dz \quad (1)$$

where AMF_G is the geometric AMF defined by the solar zenith angle (SZA) and the satellite viewing angle (VZA), $w(z)$ is the scattering weight that defines the instrument's sensitivity to NO_2 at altitude z , and $S(z)$ is a normalized vertical profile of NO_2 number density called the shape factor (Palmer et al., 2001). Scattering weights are calculated with a radiative transfer model and typically increase with altitude (Martin et al., 2002). The shape factor is usually estimated with a CTM.

Here we use scattering weights at 448 nm calculated for the application to a 435 – 461 nm spectral window compiled as a look-up table dependent on SZA, VZA, relative azimuth angle (RAA), surface albedo, cloud top pressure, and effective cloud fraction (R. Park and Kwon, 2020). We specify the shape factor with local NO_2 concentrations from the GEOS-Chem simulation described in Section 2.3. In simulations of observations from the KORUS-AQ aircraft campaign over Seoul, Yang et al. (2023) showed that GEOS-Chem was successful in reproducing the observed diurnal variation in the vertical profile of NO_2 as driven by PBL mixing, with important implications for the diurnal variation in the AMF. The GEOS-Chem simulation extends vertically to the mesosphere, and we compute in this manner the AMF for the whole atmosphere to obtain a total vertical column density. The contribution of the stratosphere to the total NO_2 column in GEOS-Chem is only 9- 20% for the Seoul and Beijing metropolitan areas of interest here. The PBL extending to 2 km altitude accounts for 95% of the tropospheric column over the Seoul Metropolitan Area (Yang et al., 2023).

2.2 Pandora data

The Pandora instruments measure radiance at 280 – 525 nm (Herman et al., 2018) and fit total column NO_2 (including the stratosphere). There are two Pandora sites in Seoul and one in Beijing (40.0°N, 116.4°E; O. Liu et al., 2023) for the 2021-2022 period. The two Pandora sites in Seoul are at Seoul National University (Seoul – SNU; 37.5°N, 127.0°E) in the southern part of Seoul (M. Kim et al., 2021; S. Park et al., 2018) and at Yonsei University (Seoul – YSU; 37.6°N, 126.9°E) in the northern part of Seoul (J. Kim., 2017). We obtain the Pandora direct Sun data from the Pandonia global network (PGN, 2023) and exclude data with a NO_2 flag of 12.

2.3 GEOS-Chem model

140 We use GEOS-Chem CTM version 13.3.4 (<https://doi.org/10.5281/zenodo.5764874>) driven by assimilated meteorological data from the Goddard Earth Observing System – Forward Processing (GEOS-FP) with a horizontal resolution of $0.25^\circ \times 0.3125^\circ$ ($\approx 25 \times 25$ km²) over East Asia ($24 - 52^\circ\text{N}$, $104 - 133^\circ\text{E}$) and 3-hourly boundary conditions from a global GEOS-Chem simulation with $4^\circ \times 5^\circ$ resolution. The model has 47 vertical levels including 14 vertical levels in the lower 2 km. Simulations were
145 conducted for DJF 2021/2022 and JJA 2022 with 6 months of initialization for each period.

Aside from emissions (see below), the simulation is the same as previously described by Yang et al. (2023) and features some modifications to the standard GEOS-Chem 13.3.4 to better reproduce KORUS-AQ aircraft observations over Korea in May-June 2016. These include aerosol nitrate photolysis, volatile chemical product (VCP) emissions and chemistry, and reduced HO₂ uptake by aerosol. Yang et al. (2023)
150 showed that the model could successfully simulate NO_x chemistry during KORUS-AQ including the diurnal variation of NO₂ vertical profiles affecting the diurnal variation of the AMF. This gives us confidence in the application of GEOS-Chem normalized vertical profiles in Eq. (1) to compute AMFs for application to the GEMS data.

Simulations for 2022 require adjustment to NO_x emissions beyond the most recent emission
155 inventories used in GEOS-Chem for China (MEIC for 2019; Zheng et al., 2021) and Korea (KORUSv5 for 2015; Woo et al., 2020). We apply for this purpose the surface NO₂ concentration trends for China from the Ministry of Ecology and Environment (MEE) network (MEE, 2023) and for South Korea from the AirKorea network (KEC, 2023). Mean 2022/2019 surface NO₂ concentration ratios in China are 0.91 in DJF and 0.83 in JJA, and mean 2022/2015 values in Korea are 0.70 in DJF and 0.51 JJA, which are applied
160 to scale the anthropogenic NO_x emissions.

Several previous studies aside from Yang et al. (2023) have evaluated the GEOS-Chem simulation of NO_x over East Asia. R. Park et al. (2021) found that GEOS-Chem successfully reproduced the NO_x vertical profiles observed during KORUS-AQ. Shah et al. (2020) found a good simulation of the

seasonality of OMI NO₂ over China and its long-term trend. M. Liu et al. (2018) found that NO₂ diurnal
165 variability at the MEE stations was well captured but the model was too low, as would be expected from
the urban nature of the sites.

2.4 Diurnal variation of NO_x emissions

Figure 1 shows the diurnal cycle of NO_x emissions used by GEOS-Chem in Beijing and Seoul.
MEIC for China provides monthly NO_x emissions separately for the transportation, residential, industrial,
170 and power sectors while KORUSv5 separates mobile, area, and point sources. Neither inventory specifies
diurnal variations in emissions. In our work, we apply the diurnal pattern from X. Liu et al. (2019) for the
power sector and Miao et al. (2020) for other sources in the MEIC inventory. For KORUSv5 we apply the
diurnal pattern from X. Liu et al. (2019) for point sources, supported by results from Bae et al. (2021), and
the industrial daily pattern from Miao et al. (2020) for area sources. We estimate the diurnal variation of
175 mobile sources in KORUSv5 using hourly Seoul Transport Operation and Information Services (TOPIS,
2023) data on weekday total traffic and construction equipment activity.

Figure 1 shows that emissions are dominated by industrial and transport sources in Beijing, and by
mobile (transport) sources in Seoul. Both sectors show a broad maximum between 7 and 18 LT that defines
the overall diurnal cycle of emissions and is similar in winter and summer. There are no significant rush
180 hour peaks in transport emissions, suggesting that the surface NO₂ maxima observed in early morning and
evening are driven more by shallow mixing depths (Moutinho et al., 2020). Total NO_x emission in Beijing
in winter is 30% greater than in summer, driven by the industrial source and possibly due to workplace
heating. There is less seasonal variation in Seoul where mobile sources are the largest emitters.

3. Intercomparison of total NO₂ columns

185 Figure 2 shows the total NO₂ columns over eastern China and South Korea observed by GEMS and
simulated by GEOS-Chem during DJF 2021/22 and JJA 2022. The GEMS data are mapped on the 0.25° ×
0.3125° GEOS-Chem grid. The yellow box delineates the Seoul Metropolitan Area (SMA). The zoomed-in
black boxes are Beijing and Seoul and the white boxes are the city centers where Pandora stations are

located. The NO₂ plume location in winter is shifted south due to the prevailing winds and the long NO_x lifetime (Seo et al, 2021). We see from Figure 2 that GEMS and GEOS-Chem have consistent spatial distributions and backgrounds, but GEOS-Chem over polluted regions is generally higher than GEMS except for Seoul in winter.

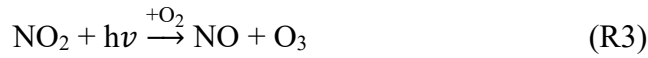
Figure 3 further intercompares GEOS-Chem and GEMS using the Pandora stations in Beijing and Seoul as evaluation metric. GEMS and GEOS-Chem reproduce the diurnal and day-to-day variability observed by Pandora in DJF ($R^2 = 0.87-0.90$) and JJA ($R^2 = 0.77-0.79$). NO₂ column magnitudes also agree well with Pandora in winter, with linear regression slopes of 0.94 for GEMS and 0.90 for GEOS-Chem. Summer shows larger biases reflecting differences between the SNU and YSU Pandora sites that cannot be resolved at the $0.25^\circ \times 0.3125^\circ$ resolution of GEOS-Chem (there are few observations at the Beijing site in JJA). YSU is more polluted than SNU, which is in a mountainous area more remote from emissions. Overall, comparison to Pandora supports the diurnal and day-to-day variability seen in the GEMS and GEOS-Chem data. The rest of our analysis focuses on the diurnal variability.

4. Diurnal variation of NO₂ columns on the urban scale

We start with an urban core analysis focusing on the white boxes shown in Figure 2 for Beijing and Seoul, representing single $0.25^\circ \times 0.3125^\circ$ GEOS-Chem grid cells where the Pandora stations are located. Scatterplot comparisons between GEOS-Chem, GEMS, and Pandora for these grid cells were shown in Figure 3. Figure 4 shows the diurnal variation of the total NO₂ column observed from GEMS and Pandora and simulated by GEOS-Chem in Beijing. GEOS-Chem results are shown as averages for all days and for the subset of days when GEMS observations are available (generally limited by cloud cover). Wintertime NO₂ in all three datasets is flat from 10 to 11 LT and then increases from 11 to 14 LT. Summertime NO₂ decreases from 8 LT to a minimum at 13-14 LT and then increases to 16 LT, consistent between GEOS-Chem and GEMS. Pandora observations in the summertime are too limited to show.

We used the GEOS-Chem budget tendency diagnostic to understand the drivers of the diurnal variation in NO₂ columns. This diagnostic tracks the mean mass-weighted changes of column

concentrations after each model operation for any selected horizontal domain, vertical column, and time
 215 period. We focus on the PBL column conservatively defined as extending to 3 km altitude after verifying
 that altitudes higher than 3 km make negligible contributions to diurnal changes in the total model column.
 Within the PBL column we consider the budget of NO₂ as that of NO_x (\equiv NO + NO₂ + NO₃ + 2N₂O₅ +
 HONO + HNO₄ + ClNO₂) multiplied by the local NO₂/NO_x PBL column concentration ratio. This
 eliminates from the budget the fast interconversion reactions within the NO_x family and provides a more
 220 useful budget perspective. The NO_x family is mainly contributed by NO and NO₂, and the main
 interconversion reactions defining the NO₂/NO_x ratio are



225 The differential net budget tendency for the PBL NO₂ column (Ω_{NO_2}) can then be related to that of NO_x
 (Ω_{NO_x}) as

$$\left[\frac{\partial \Omega_{\text{NO}_2}}{\partial t} \right]_{\text{net}} = \alpha(t) \left[\frac{\partial \Omega_{\text{NO}_x}}{\partial t} \right]_{\text{net}} \quad (2)$$

with

$$\left[\frac{\partial \Omega_{\text{NO}_x}}{\partial t} \right]_{\text{net}} = \left[\frac{\partial \Omega_{\text{NO}_x}}{\partial t} \right]_{\text{emission}} + \left[\frac{\partial \Omega_{\text{NO}_x}}{\partial t} \right]_{\text{chemistry}} + \left[\frac{\partial \Omega_{\text{NO}_x}}{\partial t} \right]_{\text{transport}} \quad (3)$$

230 and where $\alpha(t) = \Omega_{\text{NO}_2}/\Omega_{\text{NO}_x}$ is the NO₂/NO_x PBL column ratio. The terms on the right-hand side of Eq. (3)
 are updated by GEOS-Chem over its operator splitting time steps and are archived in the budget diagnostic
 as spatial and temporal averages. NO_x dry deposition is included in the emission operator, but its
 contribution is very small (Shah et al., 2020).

The second row of Figure 4 shows the different NO_x budget terms from Eq. (3) over hourly time
 235 steps, with the net tendency as the left-hand-side term. The third row shows the NO₂/NO_x PBL column
 molar ratios in GEOS-Chem. The GEOS-Chem diurnal variation in the NO₂ column in the first row reflects
 the net NO_x tendency combined with the NO₂/NO_x ratio. In this manner, we see that the increase in the NO₂
 column over the course of the day in winter reflects the dominant effect of daytime emissions, four times

higher than at night and leading to NO_x accumulation. Chemical loss is slow in winter and transport (flux
240 divergence) is the main loss term. The flat trend of the NO_2 column from 10 to 11 LT corresponds to the
diurnal minimum of the NO_2/NO_x ratio. This ratio increases over the rest of the day as the mixed layer
deepens and the freshly emitted NO is exposed to higher ozone concentrations. The increase in the ratio
contributes to the increase in the NO_2 column. The NO_2 column in GEOS-Chem thus peaks at 18 LT.
During the night, the NO_x emission decreases and the loss from transport leads to decrease in the total NO_2
245 column. The NO_2/NO_x ratio at night is only $0.55 \text{ mol mol}^{-1}$, despite no NO_2 photolysis, because of
sustained NO emission and the slow rate of the $\text{NO} + \text{O}_3$ reaction (low ozone and low temperatures).

The opposite diurnal variation of NO_2 in summer reflects weaker daytime emission of NO_x and
stronger chemical loss as shown by the GEOS-Chem budget analysis. Even though the emission term
remains larger than the chemical loss term, there is also a negative transport term because upwind
250 emissions are much lower. Without the transport loss term, the NO_2 column in summer would still increase
over the course of the day. The chemical loss of NO_x peaks at 11-12 LT and then weakens, reflecting the
noon maximum of OH concentrations combined with the decreasing NO_2 concentration, and explaining the
slow recovery of the NO_2 column in the afternoon. The NO_2/NO_x ratio is higher in summer than in winter
and shows little variation during the daytime, reflecting the higher concentrations of O_3 and HO_2 radicals
255 offsetting the effect of NO_2 photolysis. The daytime NO_2/NO_x ratio averages $0.75 \text{ mol mol}^{-1}$ in summer, as
compared to 0.5 mol mol^{-1} in winter, contributing to the seasonality of NO_2 seen from space.

Figure 5 shows the same as Figure 4 but for Seoul. The two Pandora stations show differences in
 NO_2 columns, particularly in summer, as previously shown in Figure 2. They also show some differences in
diurnal variation, particularly in winter, which we similarly attribute to local effects such as different
260 emissions, wind speeds, and geography that cannot be resolved at 25-km resolution. The diurnal variations
of GEMS and GEOS-Chem agree to within the ranges defined by the Pandora data. NO_2 columns in winter
increase from 10 to 12 LT as in Beijing but then flatten in the afternoon, which we attribute in GEOS-
Chem to stronger winds. NO_2 columns in summer show an increase from 8 to 10 LT, unlike in Beijing,

because of larger emissions initially overwhelming the chemical loss term. There follows a decrease until
265 13-14 LT and a recovery in the later afternoon, similar to Beijing and driven by the same factors.

5. Separating the influences of emission, chemistry, and transport

We showed in Section 4 that the diurnal variation of the NO₂ column observed by GEMS on the
urban scale reflects a balance between emission and transport in winter, and the added influence from
270 chemical loss in summer. This suggests that emission and chemical loss can be independently inferred from
the GEMS observations in winter and summer if the role of transport can be quantified. The transport term
can be represented with a CTM in an inversion framework (Cooper et al., 2017), but simple quantification
of the transport term on the urban scale can also be done from knowledge of the wind speed with a mass
balance approach (Jacob et al., 2016).

275 Figure 6 illustrates the sensitivity of the NO₂ diurnal variation to wind speed in the wintertime
GEMS observations over Seoul when chemical loss is a negligible term. Here the data are segregated by
hourly wind speed (6 m s⁻¹) at 850 hPa from NASA GEOS-FP meteorology. The diurnal variations are very
different at high and low wind speed, and consistent between GEMS and GEOS-Chem. At high wind
speed, the NO₂ column shows little diurnal variability because emission is balanced by transport. At low
280 wind speed, NO₂ accumulates throughout the day because the transport term is weaker and does not keep
up with emissions. But the transport term is not negligible even at low wind speed; for a 3 m s⁻¹ wind, the
ventilation time scale for the 25-km urban domain is only 2 hours. This also explains why the transport
term remains important in summer (Figures 4 and 5), as the timescale for NO_x chemical loss is about 6
hours (Shah et al., 2020).

285 One can reduce the effect of transport by spatial averaging over a large domain, thereby increasing
the ventilation timescale. Figure 7 shows the average diurnal pattern of the NO₂ column observed by
GEMS and simulated by GEOS-Chem on the ≈150-km scale of the SMA (Figure 2) in winter and summer.
Again, the diurnal variations observed by GEMS and simulated by GEOS-Chem are consistent. NO₂
columns in winter increase over the course of the day in a more regular manner than on the 25-km urban

290 scale (Figure 5) because the transport loss term is steadier and responds mainly to the change in the NO₂ column. The chemical loss term is not negligible, unlike on the urban scale, because its timescale of 24 hours is comparable to that of transport.

We see from Figure 7 that the transport term can be successfully marginalized on the scale of the SMA in summer because the chemical loss term is faster. The resulting SMA diurnal pattern of the total
295 NO₂ column is consistent with that of Seoul (Fig. 5b) but with a flatter shape and the early-afternoon minimum now driven mainly by chemistry. The amplitude is greatly dampened because emissions are five times weaker when averaged over the SMA regional domain, and because the chemical loss integrates over the residence time within the domain.

6. Conclusions

300 We used the GEOS-Chem model to interpret the diurnal variation of NO₂ columns observed from the GEMS geostationary instrument and Pandora ground-based spectrometers over East Asia in December-January-February (DJF) 2021/22 and June-July-August (JJA) 2022. This was motivated by the need to understand the unique information offered by hourly geostationary satellite observations on the budget of NO_x through the contributions of emissions, chemistry, and transport to the diurnal cycle of NO₂.

305 The GEOS-Chem model used in this work had previously shown successful simulation of the NO₂ vertical profile and its diurnal variation over Seoul in the KORUS-AQ aircraft campaign, resolving the diurnal dependence of the air mass factor (AMF) that contributes to the diurnal variation of NO₂ observed from space. Here we used the diagnostic budget capability in GEOS-Chem to isolate the contributions of NO_x emissions, chemistry, transport, and the NO₂/NO_x column ratio to the diurnal cycle of NO₂ columns.
310 We also updated NO_x emissions to 2022 including their diurnal variations. NO_x emissions for Beijing and Seoul are a factor of four higher in the daytime than at night, reflecting mobile and industrial sources, and show little variation during the daytime hours. We focused on simulation of the total atmospheric NO₂ column rather than the tropospheric column, taking advantage of the stratospheric capability in GEOS-

Chem, and to avoid errors in the definition of the tropopause. Diurnal variation of the NO₂ atmospheric
315 column is mainly determined by the planetary boundary layer (PBL) up to 3 km altitude.

We investigated the diurnal variation of the NO₂ column at the 25-km urban scale over Beijing and
Seoul. GEMS, Pandora, and GEOS-Chem show similar variability and diurnal variations. NO₂ columns in
winter increase over the course of the daytime hours, reflecting accumulation from high daytime emissions
offset by loss from horizontal transport (flux divergence), and further enhanced by increase in the NO₂/NO_x
320 over the course of the day as ozone is entrained in the growing mixed layer. Chemical loss of NO_x in winter
is too slow to play a significant role in the observed diurnal variation. In summer, by contrast, NO₂ columns
decrease from 10 to 14 local time (LT) because of NO_x photochemical oxidation compounding the loss
from transport. The NO₂ column would still increase throughout the day in summer as in winter were it not
for the loss from transport.

325 Our results indicate that the diurnal variation of NO₂ column observed from geostationary orbit can
be used to quantify urban NO_x emissions in winter, and chemical loss in summer, but the transport term
must be accounted for. This can be done by simple mass balance using knowledge of the local wind speed
or by an inversion with a full CTM.

We further examined the importance of transport for interpreting the diurnal variation in the GEMS
330 urban NO₂ data by segregating the Seoul data by wind speed. In winter, the low-wind GEMS data ($< 6 \text{ m s}^{-1}$)
show steady rise of NO₂ over the course of the day while the high-wind data ($\geq 6 \text{ m s}^{-1}$) show flat diurnal
variation, consistent with the GEOS-Chem the model. Transport plays an important role in the NO_x budget
in both cases but cannot keep up with the high daytime emissions in the low-wind case.

We examined whether the role of transport in the diurnal variation of the urban NO₂ column could
335 be reduced by spatial averaging of the data over the 150-km regional scale of the Seoul Metropolitan Area
(SMA). The SMA data in winter show a steady increase over the daytime hours due to emissions, but the
transport term remains the major sink of NO_x. The SMA data in summer show negligible loss from

transport in daytime because chemical loss term is much faster, but the diurnal amplitude is weak because of diluted emissions and long residence times for the air over the regional domain.

Our conclusions regarding the interpretation of the diurnal variation of NO₂ columns observed by GEMS can be extended to other instruments of the geostationary air quality constellation, such as TEMPO over North America, launched in April 2023 and Sentinel-4 over Europe, scheduled for launch in 2024. This work further lays the groundwork for use of GEOS-Chem in inversions of the geostationary satellite data to infer NO_x emissions.

Code Availability

The model code used in this work is available at <https://doi.org/10.5281/zenodo.5764874> (The International GEOS-Chem User Community, 2021).

Data availability

The GEMS L2 NO₂ v2.0 slant column data can be obtained with a request to the NIER GEMS team (NIER, 2023). The total NO₂ columns from Pandora are available from the Pandonia Global Network website (<http://pandonia-global-network.org>; PGN, 2023). The surface NO₂ data over China are available from <https://quotsoft.net/air/> (MEE, 2023). The surface NO₂ data over South Korea are from AirKorea website (<https://www.airkorea.or.kr>; KEC, 2023). The Seoul hourly traffic count data is available from the Seoul Transport Operation and Information Services website (<https://topis.seoul.go.kr/>; TOPIS, 2023).

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Author contribution

The original draft preparation was done by LHY, with review and editing by DJJ, KRT, JHC, HL, and JK. DJJ contributed to project conceptualization. Modeling was done by LHY with additional support from HPL, NKC, SZ, VS, EB, XF, RMY, and DCP. The formal analysis was conducted by LHY with additional support from DJJ, RD, YJO, HC, JP, HLL, WJL, SK, EK, KRT, and JHC.

Competing Interests

The contact author has declared that none of the authors has any competing interests.

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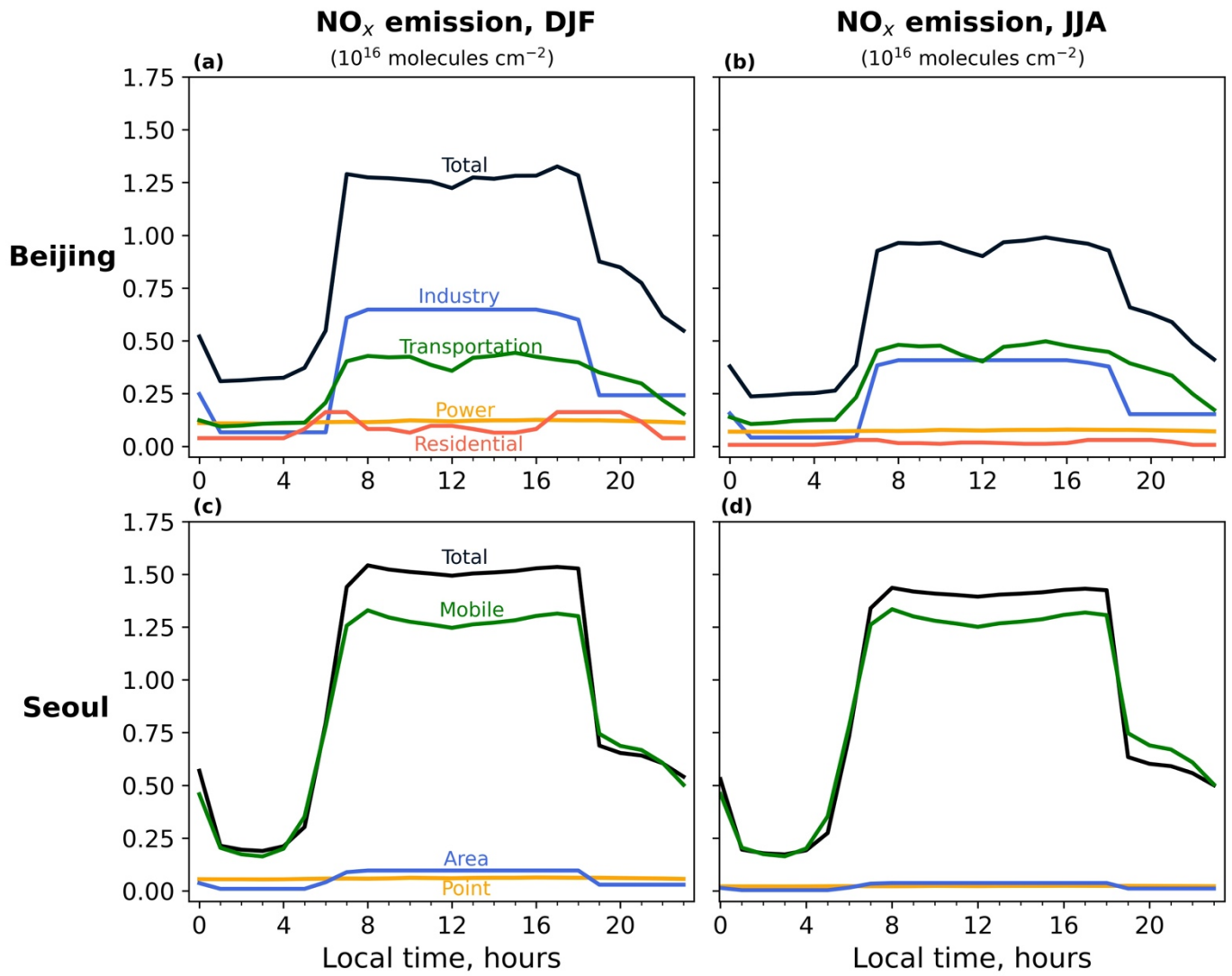


Figure 1. Diurnal variation of NO_x emissions in Beijing and Seoul for DJF 2021/2022 and JJA 2022. Local time is Chinese Standard Time (CST) for Beijing and Korean Standard Time (KST) for Seoul. Solar noon is at 12:08 – 12:27 CST in Beijing and 12:21 – 12:45 KST in Seoul. Values are for the white boxes in Figure 2. Different colors represent different sectors, and the black line shows the total emission.

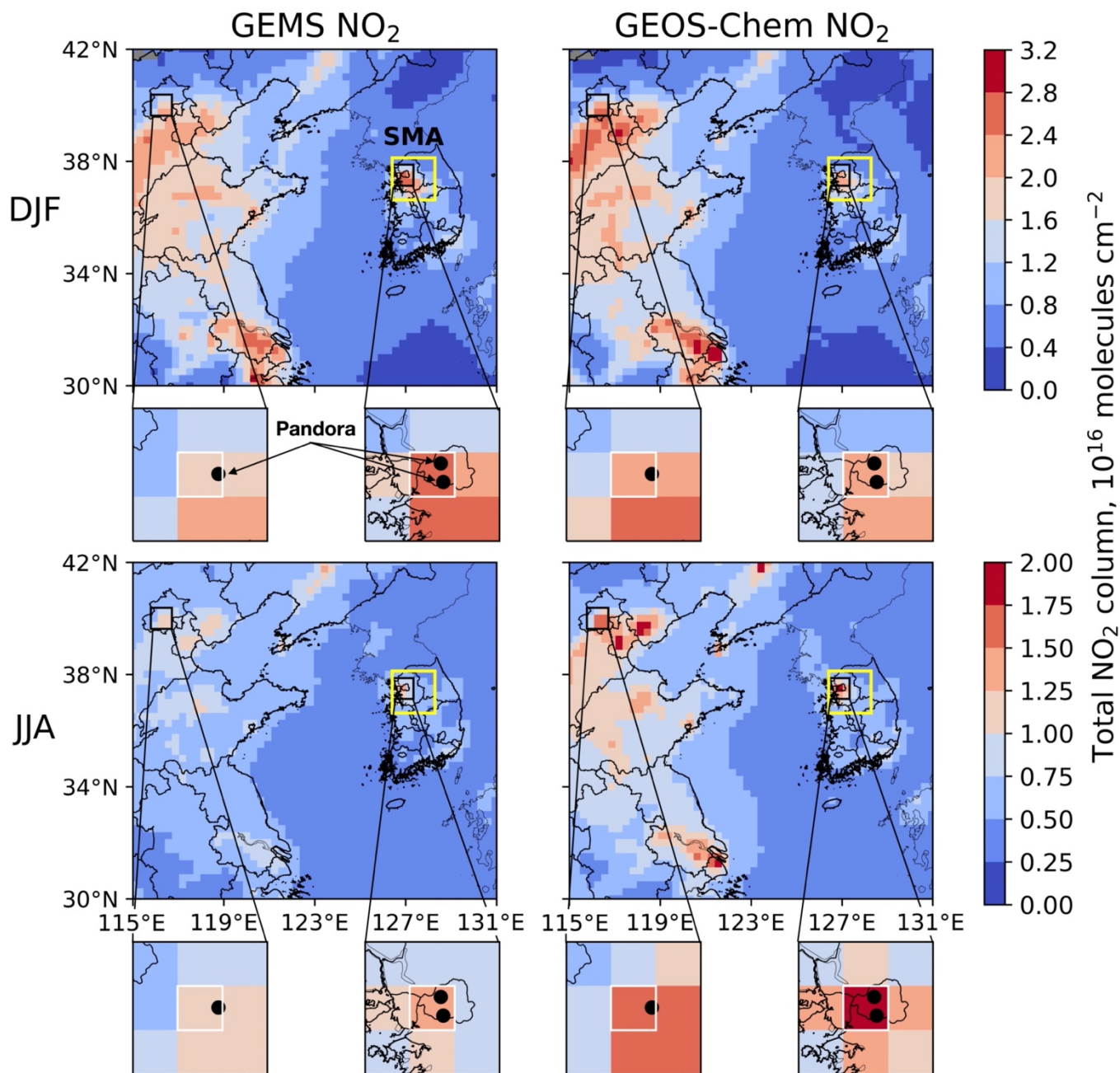


Figure 2. Total NO₂ columns over East Asia retrieved by GEMS and simulated by GEOS-Chem. The data are 3-month averages for December-January-February (DJF) 2021/22 and June-July-August (JJA) 2022 on the 0.25° × 0.3125° GEOS-Chem nested grid. The yellow rectangle delineates the Seoul Metropolitan Area (SMA; 36.6-38.1°N, 126.4-128.3°E). The zoomed-in plots show Beijing and Seoul, and the white boxes are the 0.25° × 0.3125° urban cores where the Pandora stations are located (black circles). Scales are different for DJF and JJA.

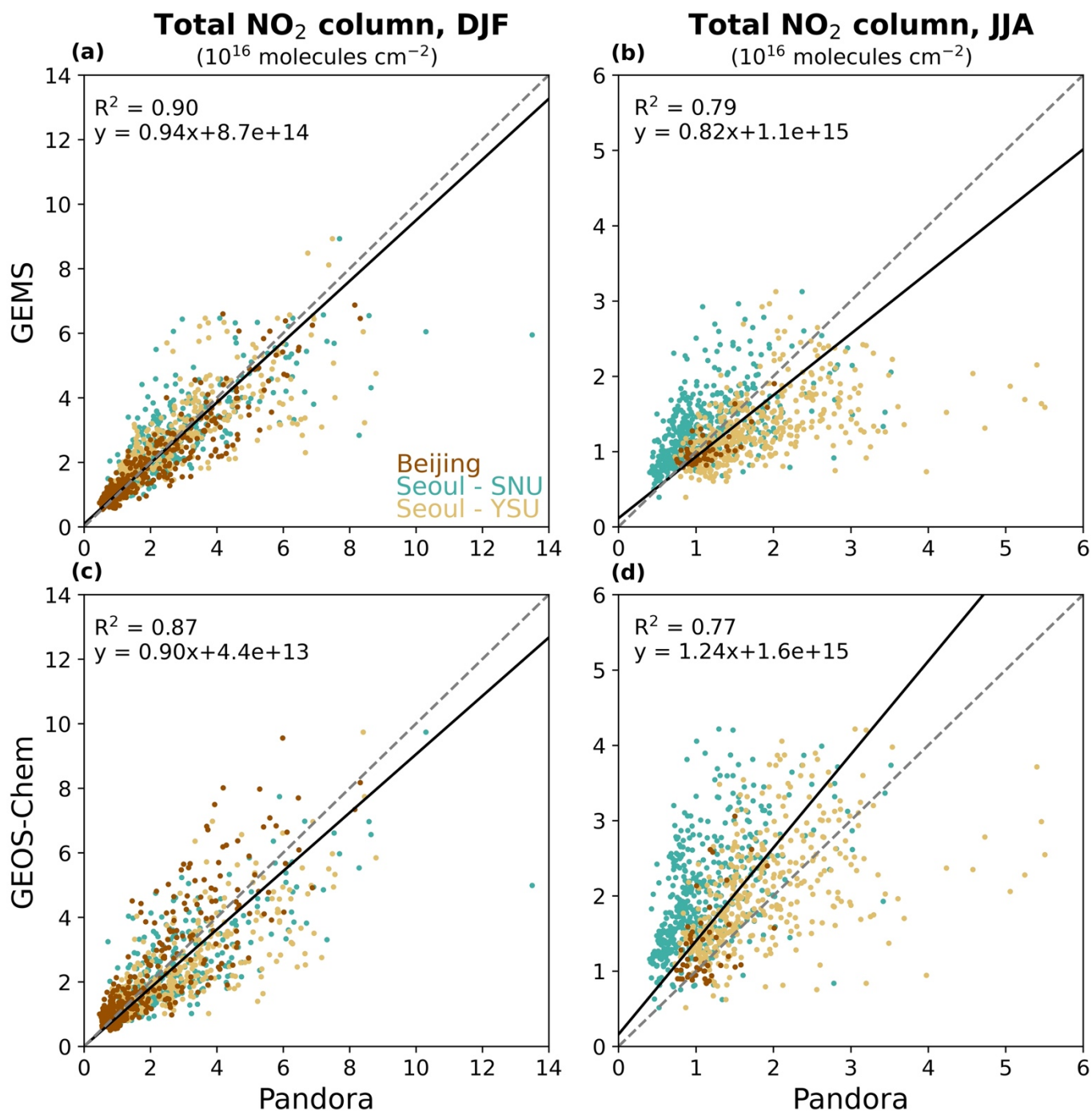


Figure 3. Intercomparison of GEOS-Chem, GEMS, and Pandora NO₂ columns for the Pandora sites in Beijing and Seoul. The Figure shows scatterplots of daytime hourly data for DJF 2021/2022 and JJA 2022. GEMS is mapped on the 0.25° × 0.3125° GEOS-Chem grid. Coefficients of determination (R^2) and reduced-major axis linear regressions are shown. The 1:1 line is dashed. The Beijing Pandora site has limited observations in JJA.

Beijing, China

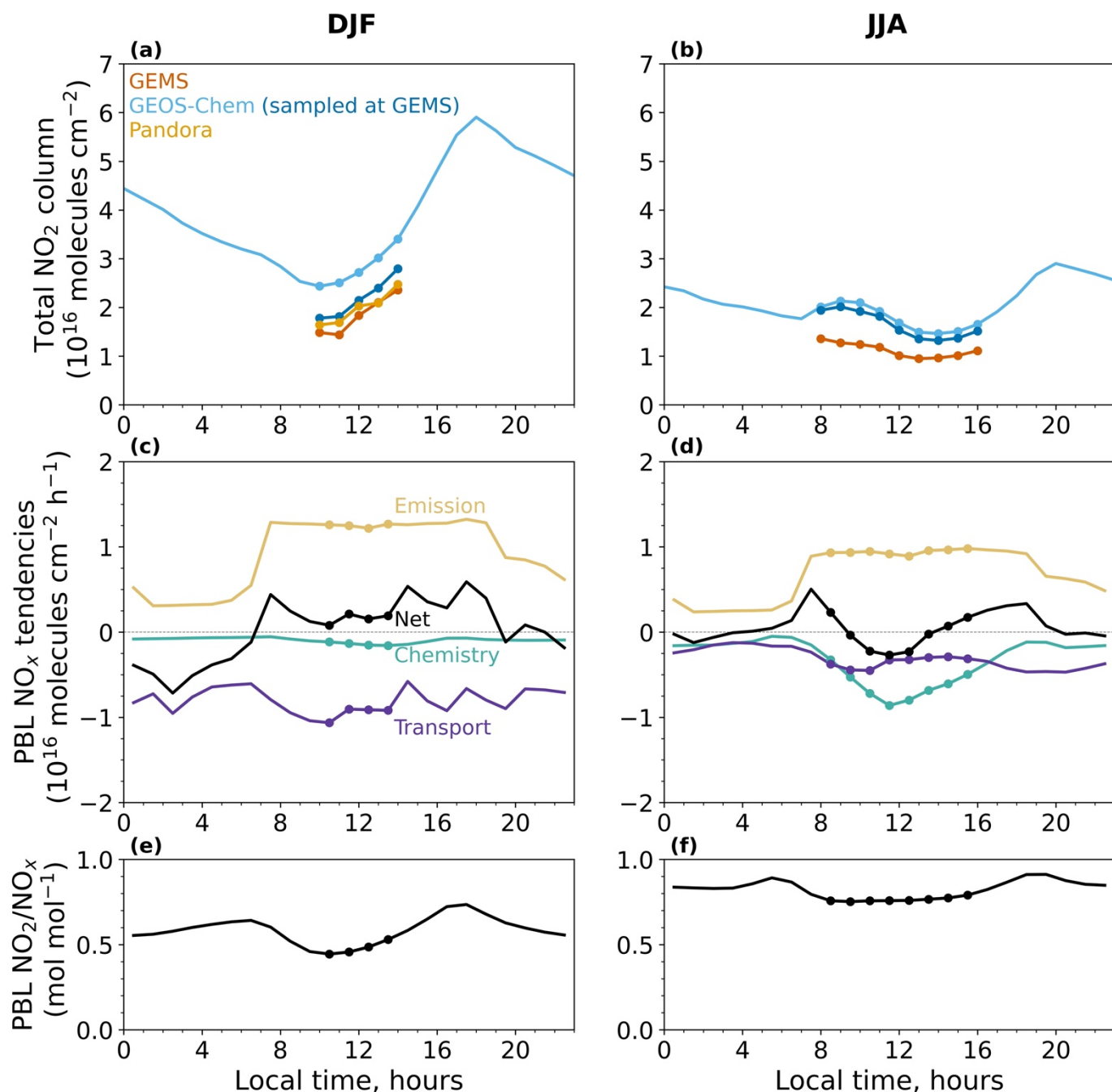


Figure 4. Diurnal variation of total NO_2 column and driving processes in Beijing. The first row shows the average NO_2 columns observed by GEMS and Pandora, and simulated by GEOS-Chem, in DJF 2021/22 and JJA 2022 for the $0.25^\circ \times 0.3125^\circ$ GEOS-Chem grid cell in the urban core where the Pandora station is located (white box in Figure 2). GEMS observations are available for the hours indicated by symbols. GEOS-Chem results for the full diurnal cycle are shown as averages for all days and for the subset of days when GEMS data are available (generally limited by cloud cover). Pandora data are not shown for JJA due to a limited number of observations (Figure 3). The second row shows the hourly tendencies in the GEOS-Chem NO_x budget (averaged for all days) for the planetary boundary layer (PBL) conservatively defined as extending up to 3 km altitude. The tendencies describe the contributions from individual processes to the NO_x budget as given by Eq. (3), with NO_x defined as $\text{NO}_x \equiv \text{NO} + \text{NO}_2 + \text{NO}_3 + 2\text{N}_2\text{O}_5 + \text{HONO} + \text{HNO}_4 + \text{ClNO}_2$. The third row shows the PBL NO_2/NO_x column molar ratio in GEOS-Chem.

Seoul, South Korea

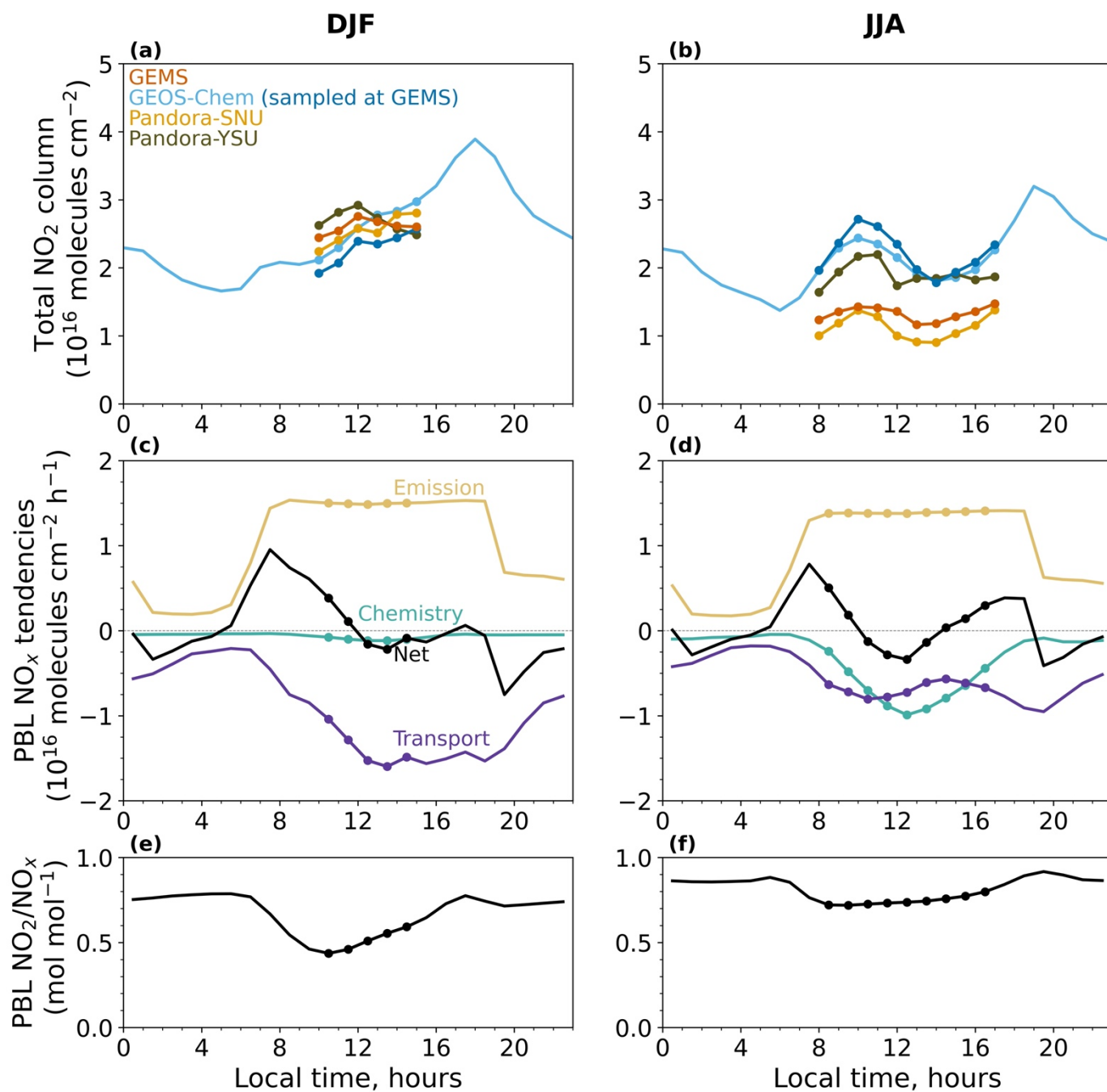
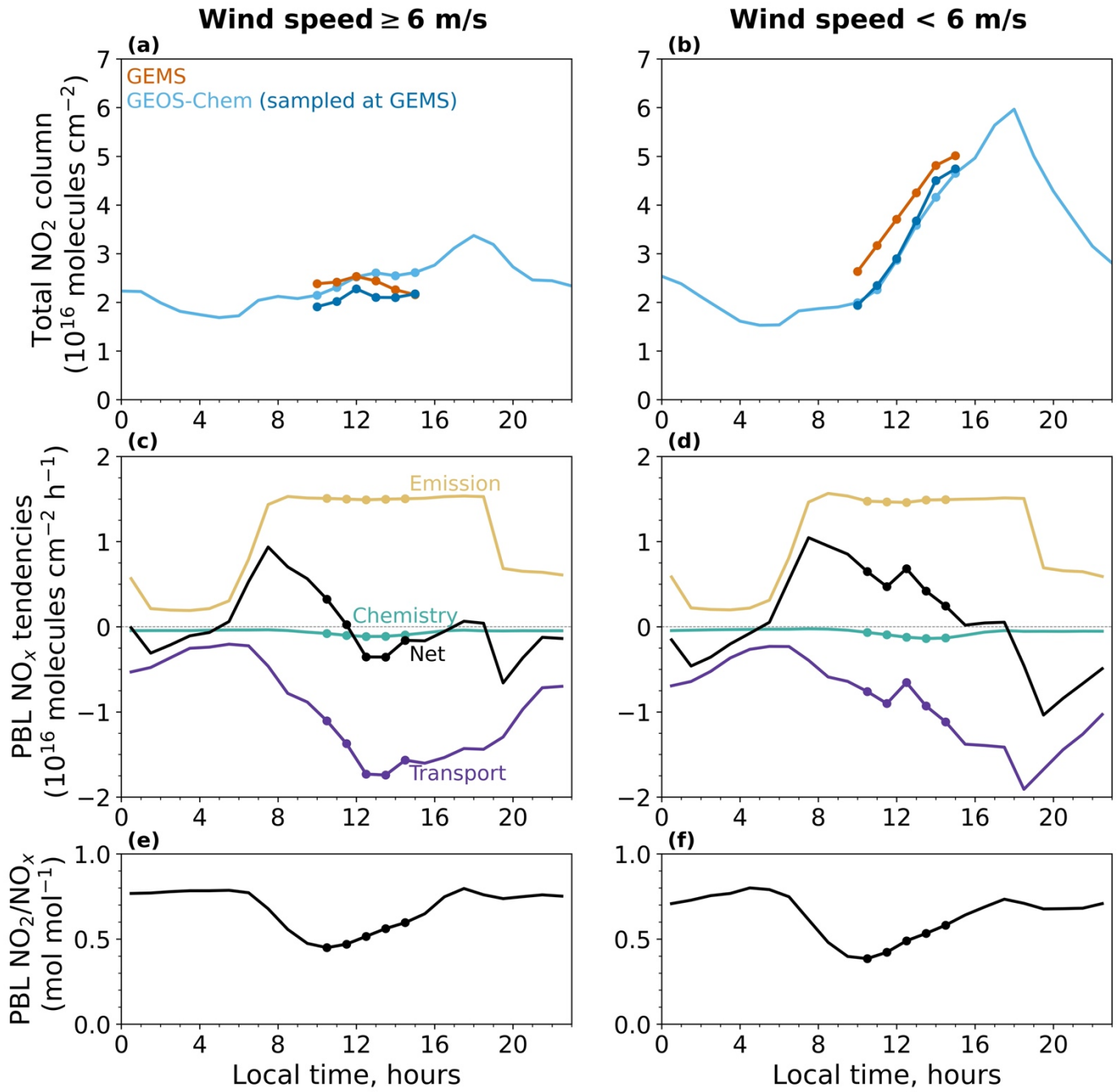


Figure 5. Same as Figure 4 but for Seoul.

DJF, Seoul



595 **Figure 6.** Same as Figure 4 but for DJF 2021/22 in Seoul with data segregated by wind speed. Segregation threshold is 6 m s^{-1} for the 850 hPa hourly wind speed in the NASA GEOS-FP meteorological data used as input to GEOS-Chem.

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SMA, South Korea

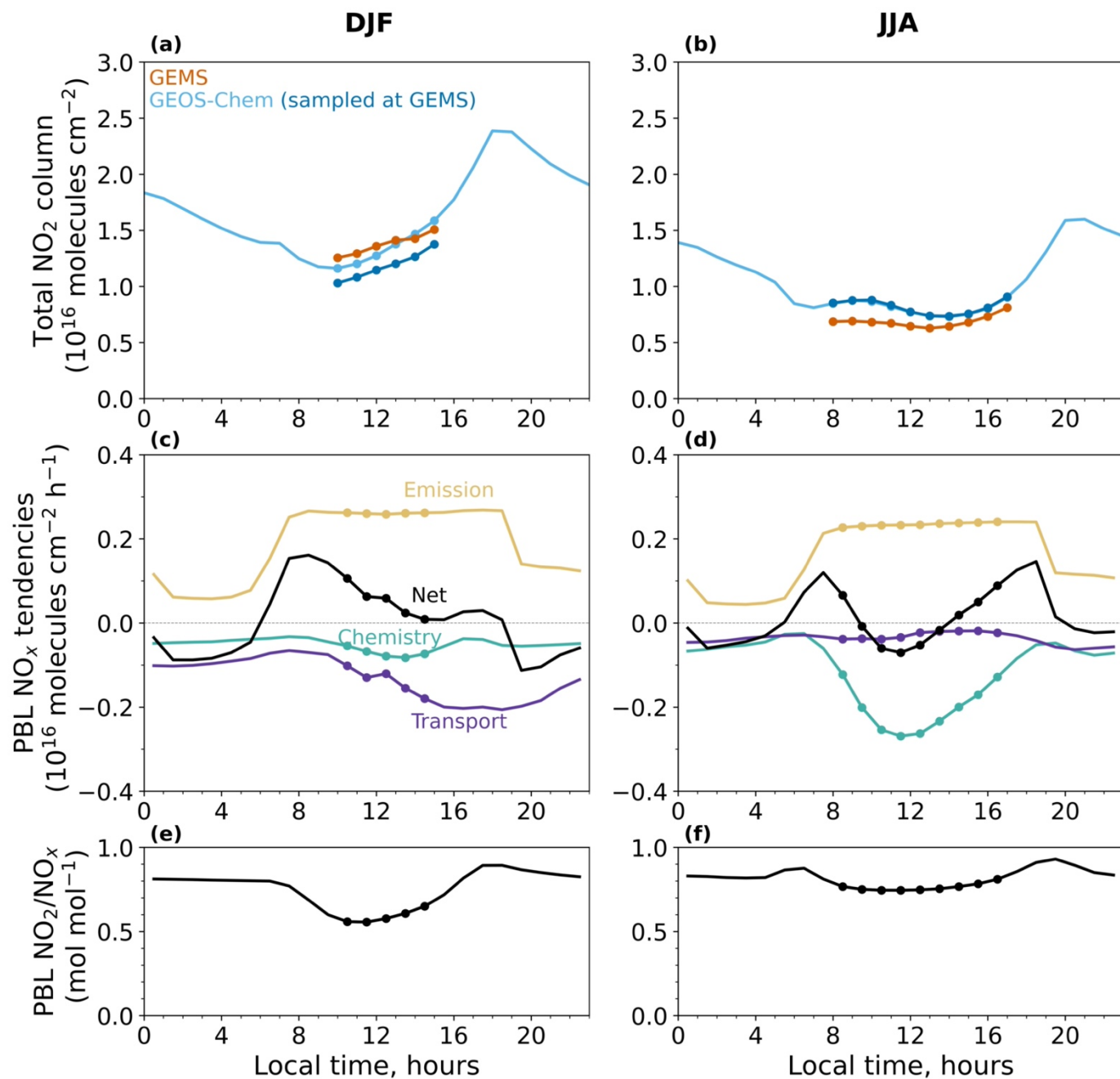


Figure 7. Same as for Figure 4 but for the Seoul Metropolitan Area (SMA; 36.6-38.1°N, 126.4-128.3°E) corresponding to the yellow box in Figure 2.